# Sixth Annual Conference on Carbon Capture & Sequestration

Geologic Storage-MMV (1)

Monitoring spatio-temporal variability of surface CO<sub>2</sub> emissions at the Horseshoe Lake tree kill, Mammoth Mountain, CA

Jennifer L. Lewicki, George E. Hilley, Marc L. Fischer, Toshi Tosha, Ryosuke Aoyagi, Koji Yamamoto, and Sally M. Benson

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Jennifer L. Lewicki, Lawrence Berkeley National Laboratory

George E. Hilley, Stanford University

Mare I. Fischer, Lawrence Berkeley, National Laboratory

Marc L. Fischer, Lawrence Berkeley National Laboratory Toshi Tosha, Geological Survey of Japan

Ryosuke Aoyagi, Mizuho Information and Research Institute

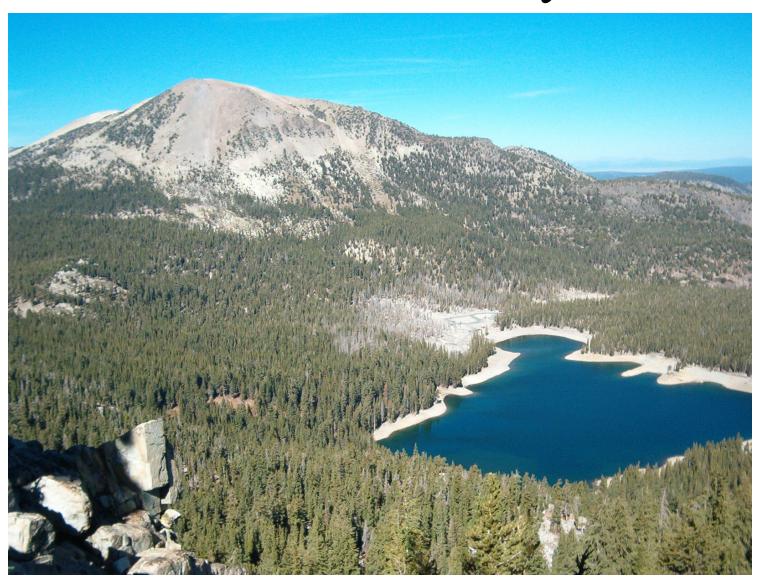
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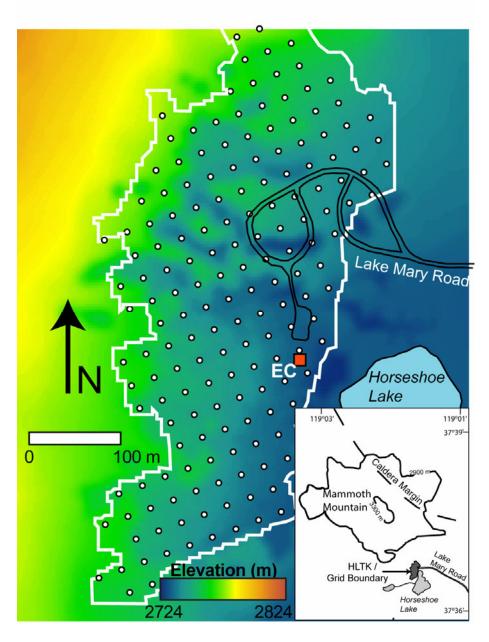
#### Motivation

- Mammoth Mountain studied as a natural analogue for large CO<sub>2</sub> leakage events
- Investigate spatio-temporal variability of soil CO<sub>2</sub> fluxes and relationship to meteorological parameters and topography using chamber method
- Compare chamber and eddy covariance CO<sub>2</sub> flux measurements at challenging site

# Recent Activity



# Study Site



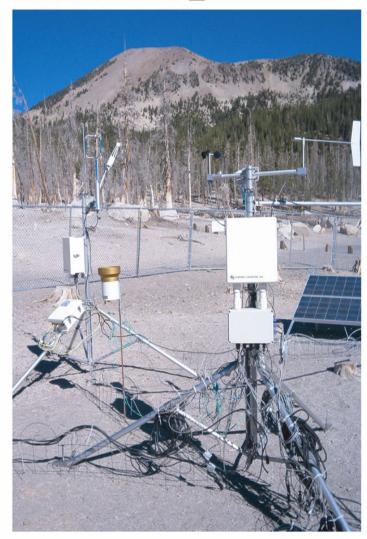
# Chamber Soil CO<sub>2</sub> Fluxes



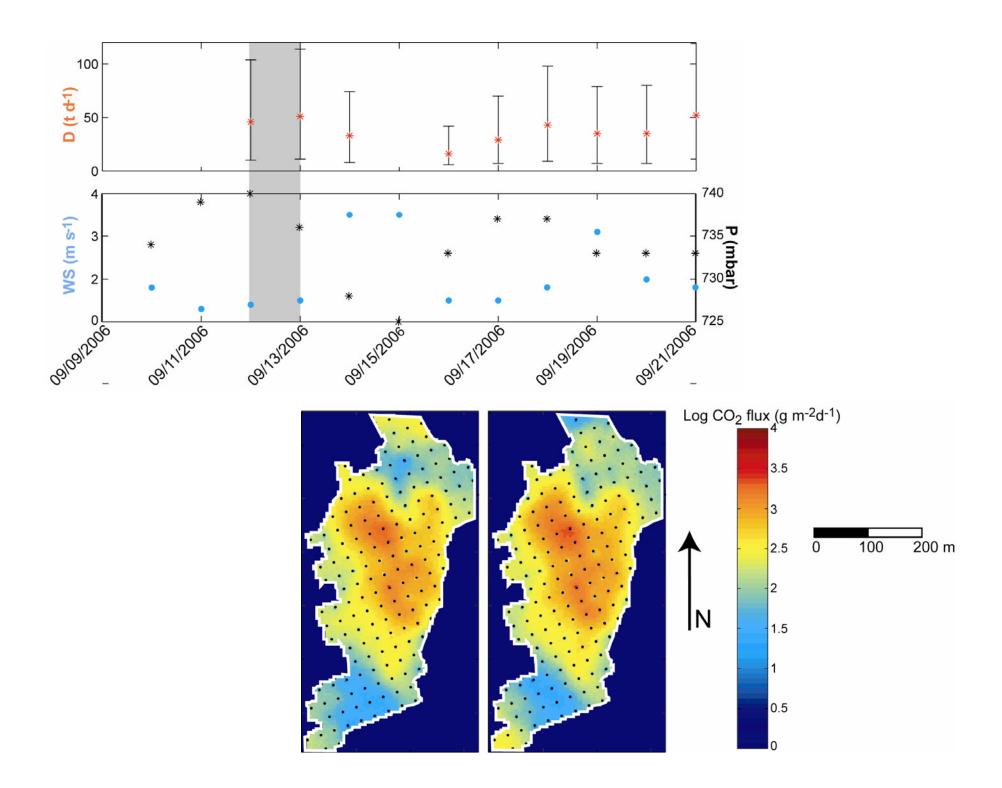
#### **Eddy Covariance Net Surface CO<sub>2</sub> Flux**

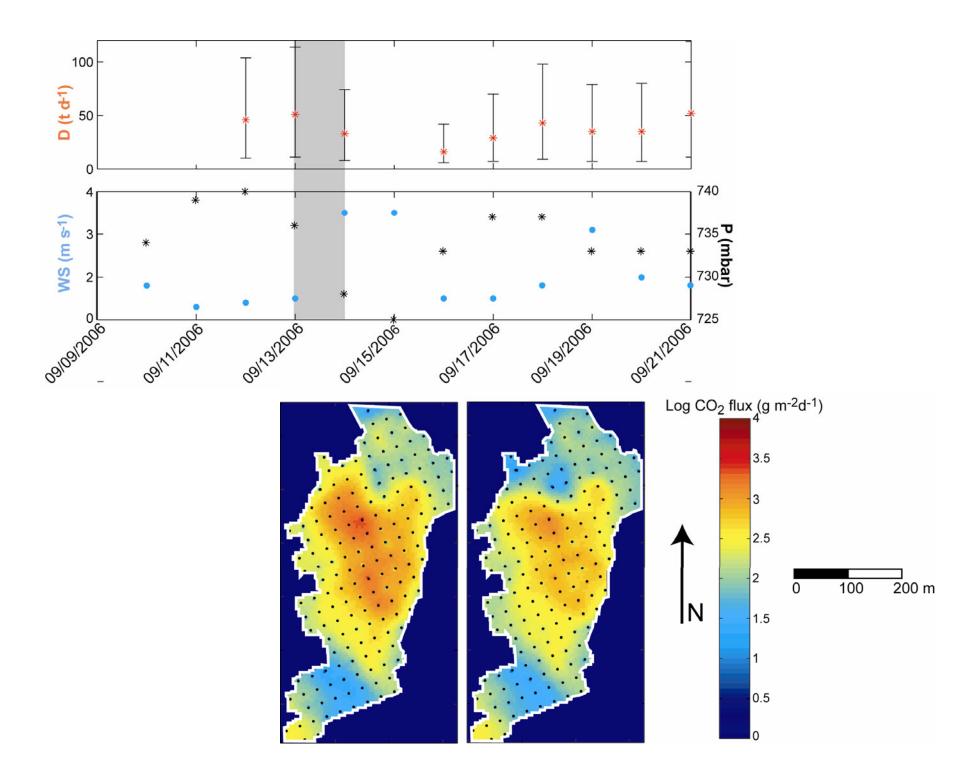
$$F_c = \overline{w'c'}$$

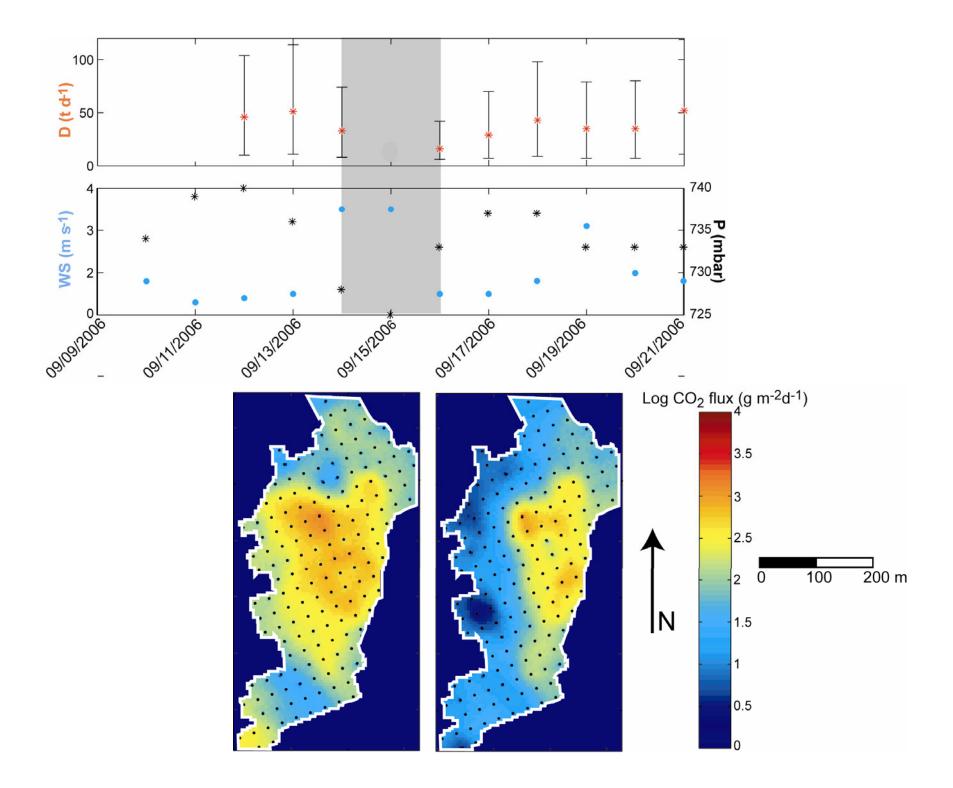
- Requires sufficiently long averaging time, steady-state conditions; assumes homogeneous surface
- Data filterered for systematic errors
- $F_c$  is integral of surface flux over upwind footprint (m²-km² scale) that scales with measurement height

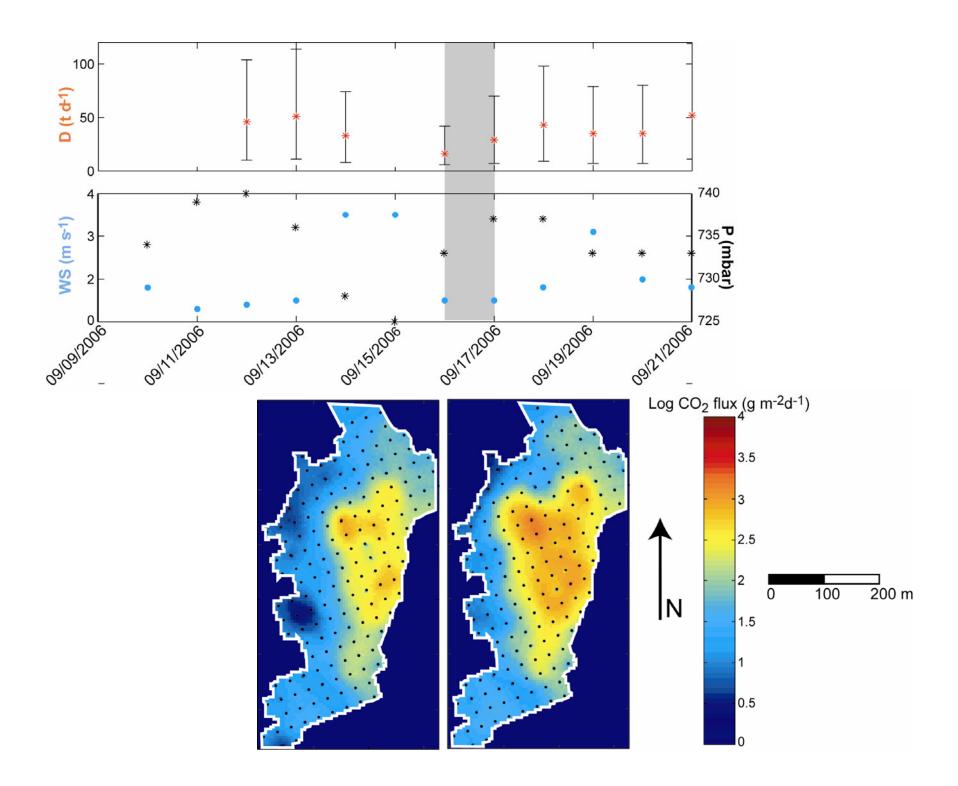


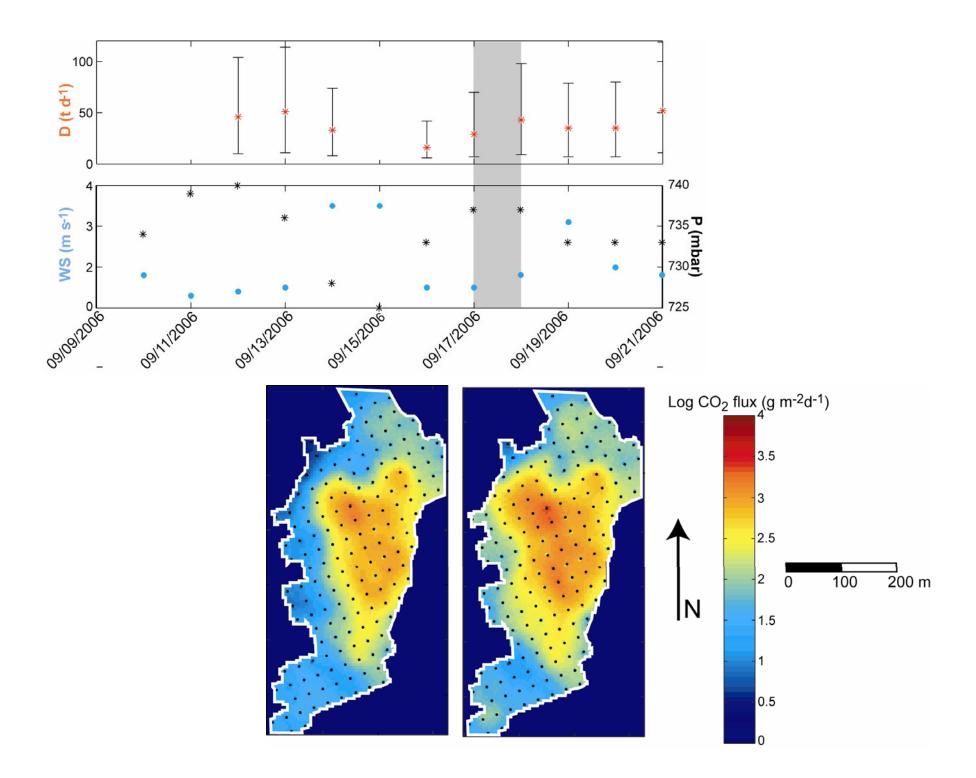
-Additional parameters: Atmospheric pressure, temperature, radiation, and humidity, soil temperature, moisture, and heat flux

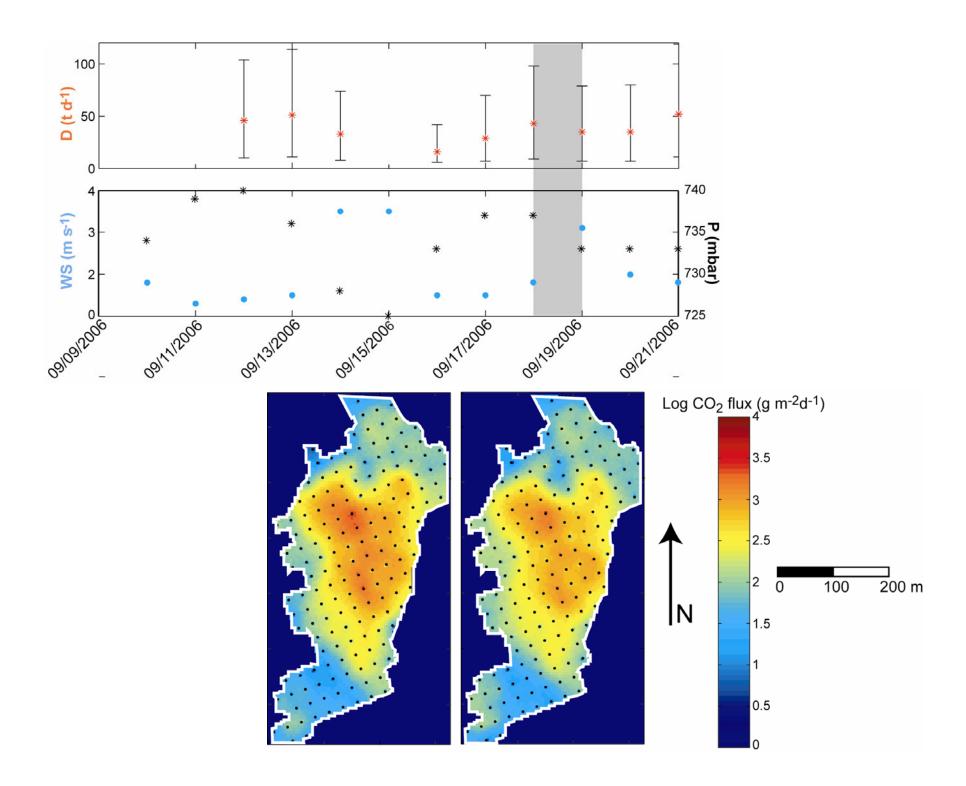


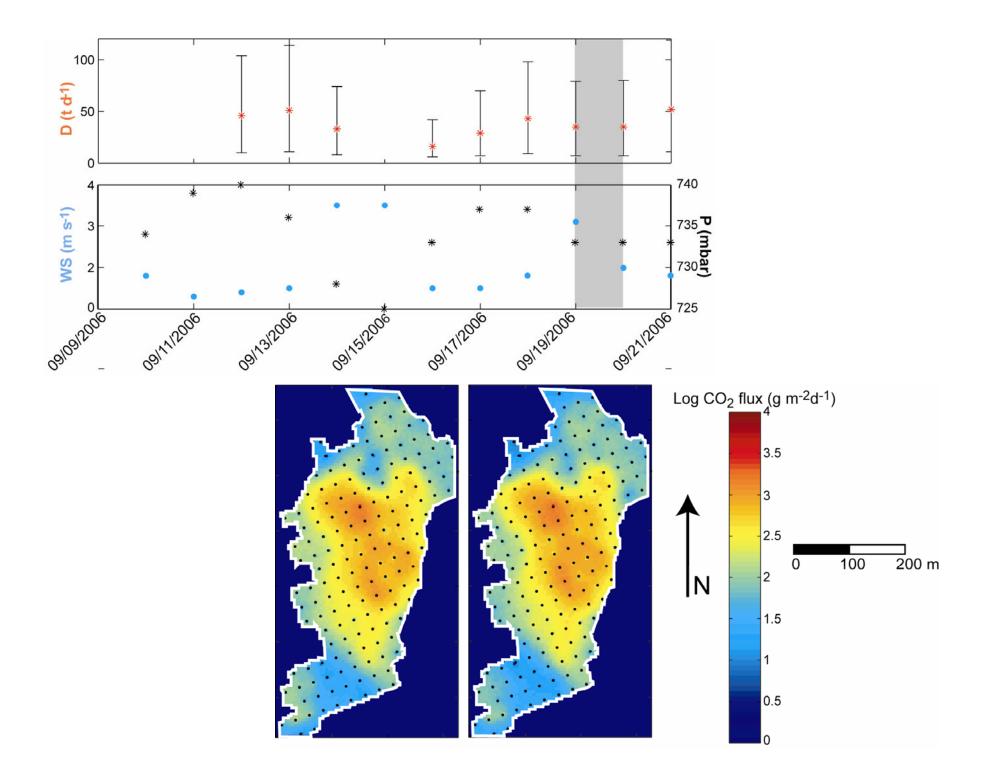


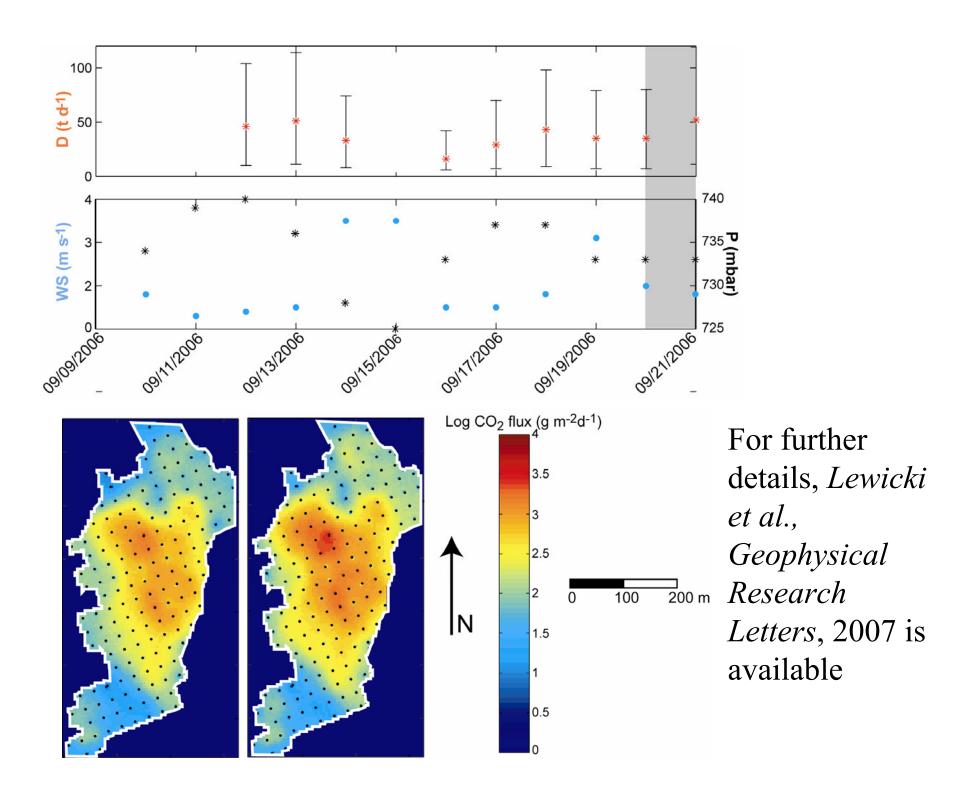




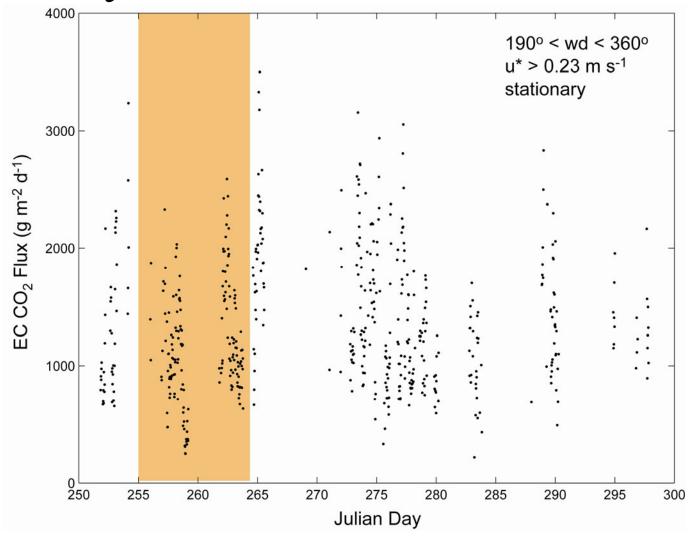




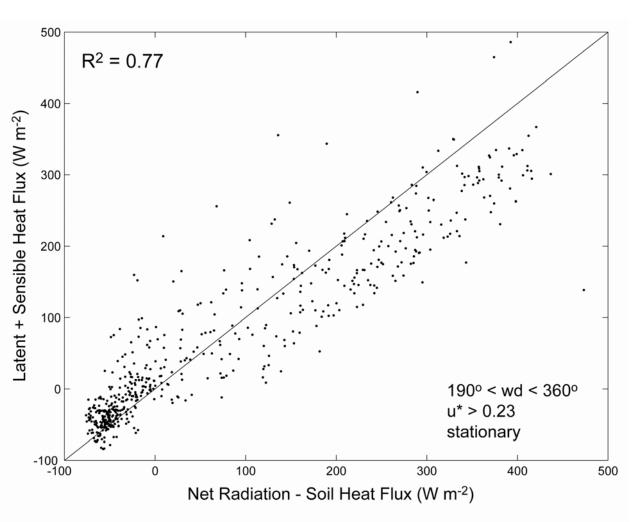




## Eddy Covariance Time Series



#### Energy Balance



Latent heat flux (LE) = F<sub>H2O</sub> x latent heat of vaporization of water
- heat flux associated with evaporation (+) and condensation (-)

Sensible heat flux (H) =  $\rho c_p \overline{T'w'}$ 

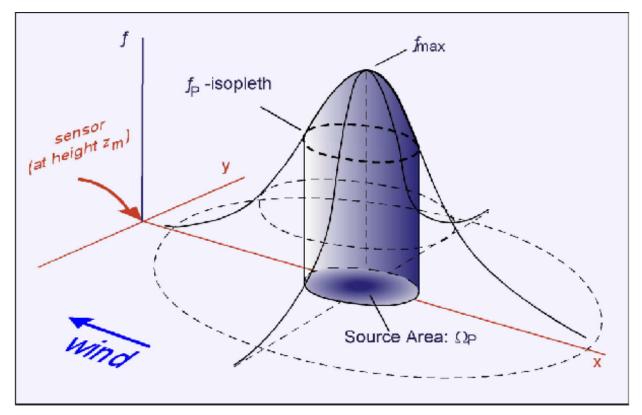
 $\rho$  = density dry air

c<sub>p</sub> = heat capacity of water at constant P

T' = fluctuating sonic temperature

w' = fluctuating vertical wind speed

- heat flux associated with temperature gradient between surface and atmosphere (+ when surface is rel. warm, - when atm. is rel. warm)



# EC Footprint

$$\eta(x_{m}, y_{m}, z_{m}) = \int_{-\infty-\infty}^{\infty} Q_{\eta}(x', y', z' = z_{0}) \cdot f(x_{m} - x', y_{m} - y', z_{m} - z_{0}) \cdot dx' dy'$$

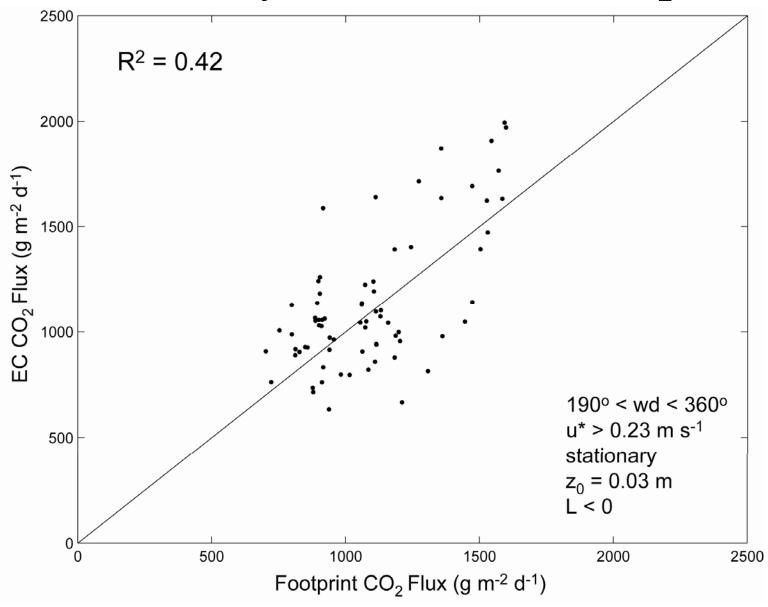
n = value of measured quantity (e.g., flux) at point  $(x_m, y_m, z_m)$  originating from the source at the surface  $(x', y', z' = z_0)$  with strength  $Q_n$ 

f= the footprint (or source weight) function, a probabilistic weighting function that assigns a relative weight to each of the source strengths  $Q_n$ , depending on separation distance between the measurement and the source

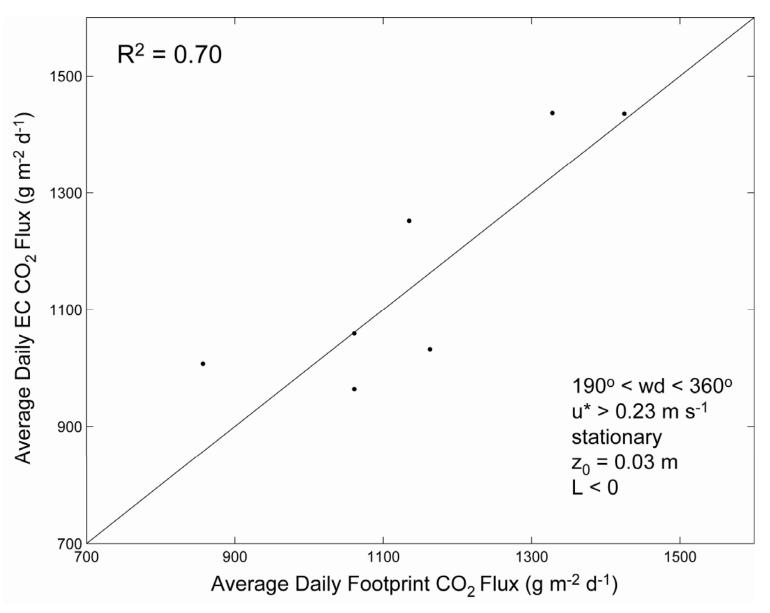
f is small for small separation distances, rises to a maximum with increasing distance, then falls off as separation is increased.

For fluxes, f is determined based on 2-D advection-diffusion equation. Depends on mean ws, direction, atmospheric stability, surf. rough.

#### Chamber-Eddy Covariance Comparison



## Average Daily Flux Comparison



#### Summary

- Large, previously undocumented, spatiotemporal variations in soil CO<sub>2</sub> flux over multiple days associated with a weather front observed by AC method
- Potential effects of topography and meteorological parameters should be considered prior to the placement of continuous monitoring devices and interpretation of time series of data

#### Summary

- EC deployed at a site that challenged the basic assumptions of the measurement
- Based on energy balance closure, EC performed well.
- Moderate to good correlation between AC and EC fluxes based on footprint modeling observed

#### Summary

- Sources of error to consider in chamber-EC comparisons:
  - Heterogeneous CO<sub>2</sub> source distribution, complex terrain introduce error into EC measurement and footprint modeling.
  - Temporally varying source flux distribution difficult to capture with chamber measurements on time scales less than inter-daily.
- EC can be used to monitor surface fluxes in challenging environments, but data gaps must be tolerated. EC best used in conjunction with complementary AC method.

#### Thank you

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